

The background of the slide is a blue-tinted photograph of a river. In the foreground, there are several reeds or grasses growing out of the water. The water surface is covered with numerous small, concentric ripples, suggesting it has recently rained. The overall scene is calm and natural.

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A STUDY OF SEDIMENTATION IN RIVERS USING SEDIMENT TRAPS

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Suspended sediment origin

- watershed-sourced sediment
(slope wash, gully erosion, volcanic eruptions)

$$d_w < d^*$$



- channel-sourced sediment

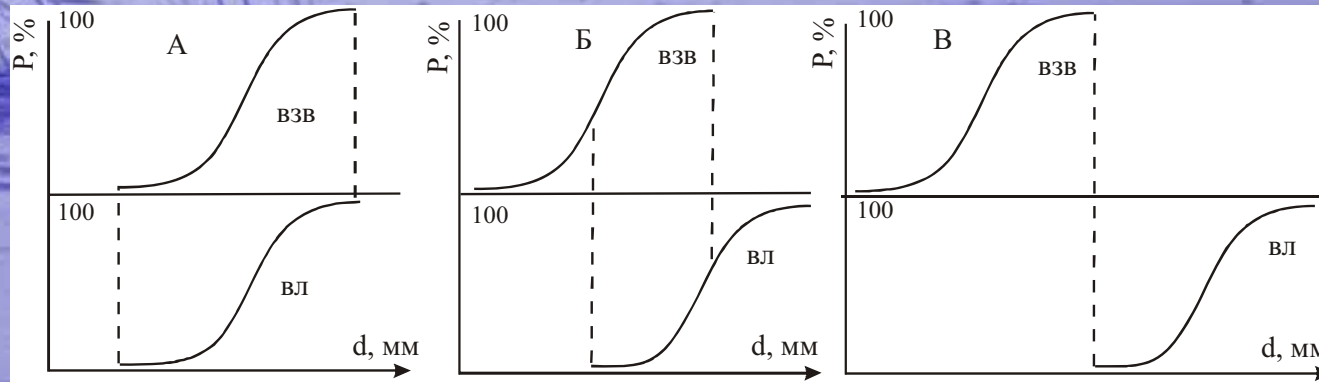
$$d_{ch} \geq d^*$$

$$d^* = 0,05 \text{ mm}$$

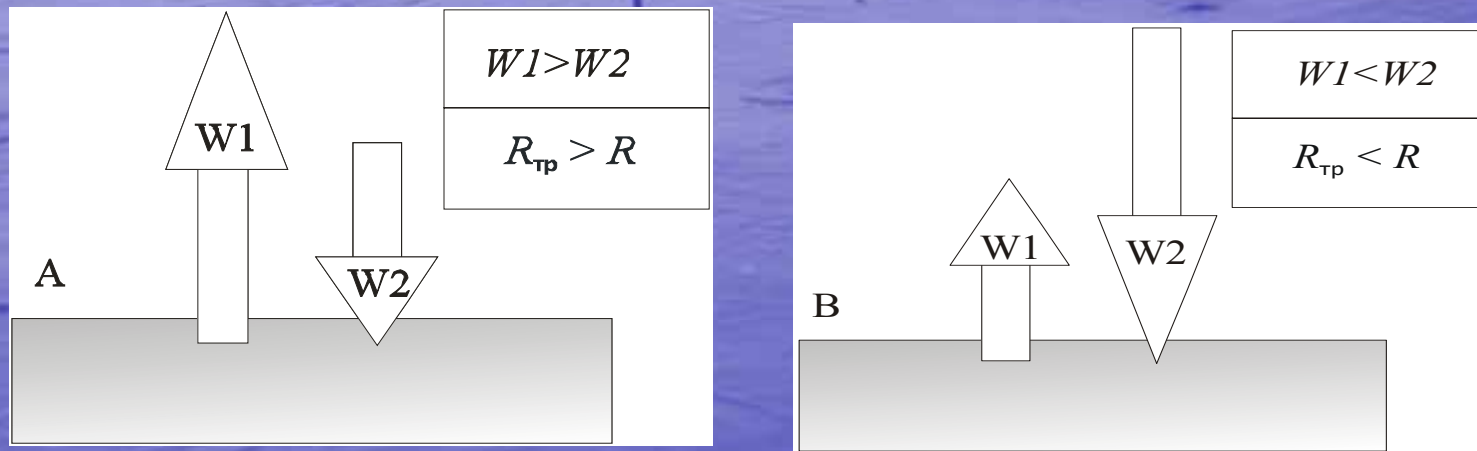


Channel-flow interaction

Comparison between bed load (вЛ) and suspended load (вЗВ)



Interaction between bed deposits (W) and suspended load (R)



$$W_2 - W_1 = \Delta W$$

W_1 – sediment input (m³), W_2 – sediment output from the channel reach (m³), ΔW – sediment budget (m³),

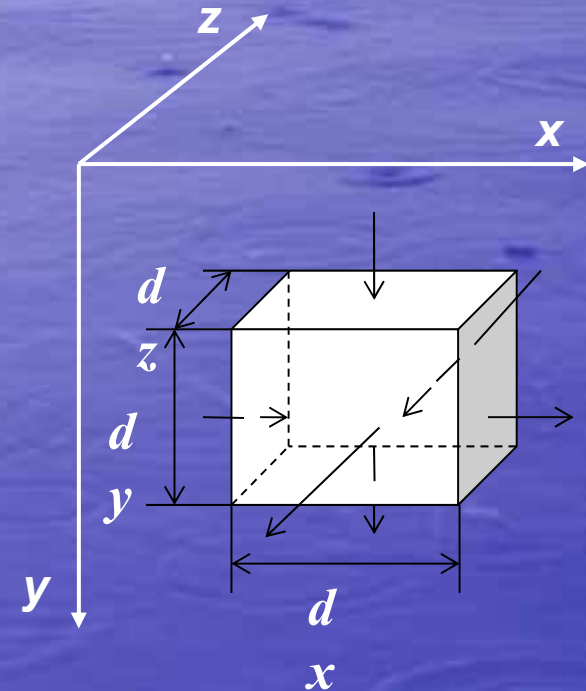
Suspended sediment concentration

$$\frac{\partial s}{\partial t} = \frac{A}{\rho} \left(\frac{\partial^2 s}{\partial x^2} + \frac{\partial^2 s}{\partial y^2} + \frac{\partial^2 s}{\partial z^2} \right) - \left(v \frac{\partial s}{\partial x} + u \frac{\partial s}{\partial y} + w \frac{\partial s}{\partial z} \right) - \omega \frac{\partial s}{\partial y}$$

- Turbulent advection – longitudinal changes of water body parameters
- Dispersion – transverse changing of parameters
- Convection – vertical changing of parameters



gravity



Methods

1. Sediment installation on the river bottom



Small trap:

2 sm diameter and 20 sm height.

Large trap:

height 30–50 sm, diameter 4,4 sm

2. Comparison with published data

Kozerski HP, Leuschner K. New plate sediment trap: design and first experiences. Verh Intern Verein Limnol 2000;

Kozerski HP. Determination of areal sedimentation rates in rivers by using plate sediment trap measurements and flow velocity—settling flux relationship. Water Res 2002;36:2983–90.

Wilson A.J., Walling D.E., Leeks G.J.L. In-channel storage of fine sediment in rivers of southwest England // Sediment transfer through fluvial system. IAHS Publ. 288. 2004. P.291-299

Walling D.E., Amos C.M. 1999. Fine sediment in a chalk stream system // Hydrological processes. Vol. 13. P. 323-340

Parameters of sediment traps for rivers

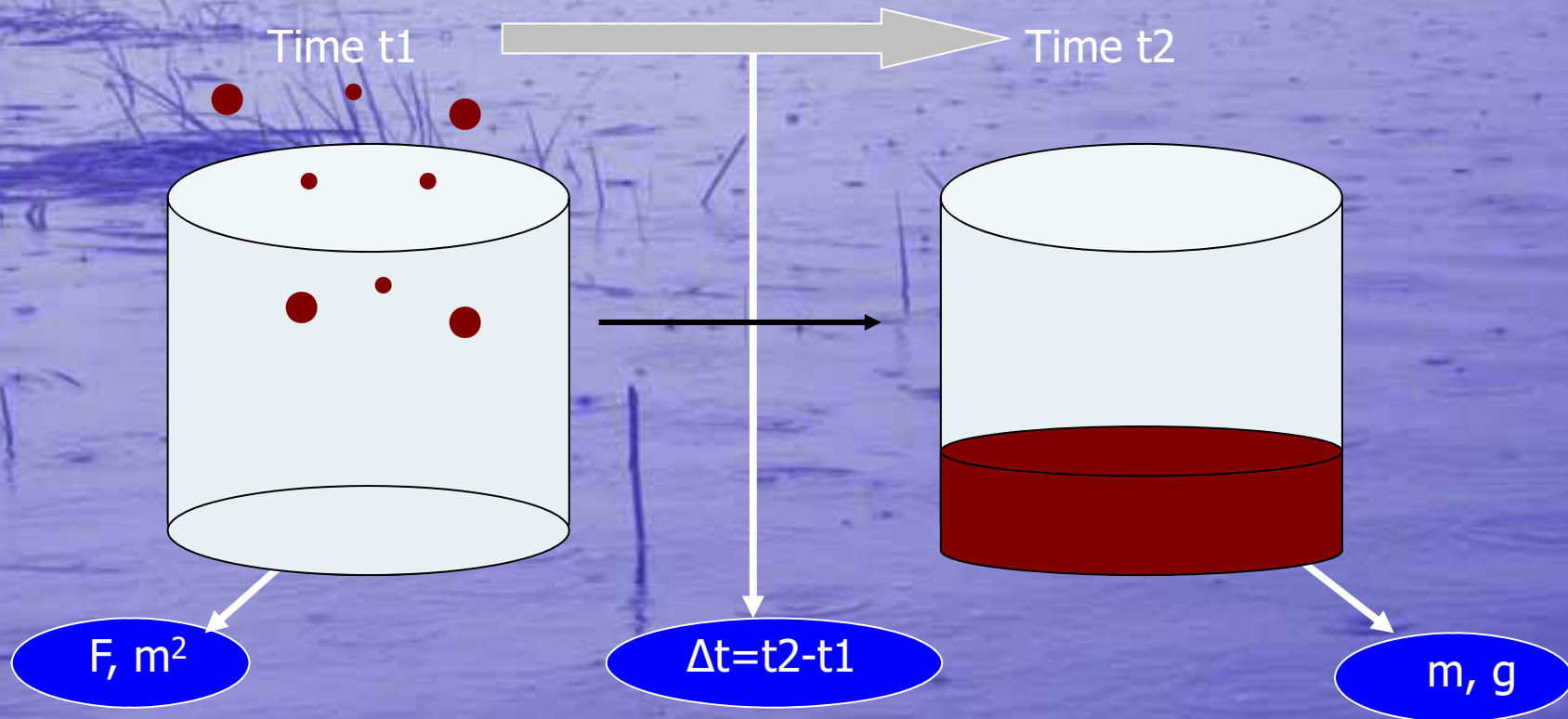
Type	Height, m	Diameter, m	Bottom area F, m ²	Volume, m ³	Total trap weight, kg	Location, m
Large	0,4	0,075	0,004	0,001	5	1 - on the bed level 2 - above bottom 3 – above bottom
Small	0,25	0,04	0,002	0,001	2	1 - on the bed level



- Field-based sediment trapping



Sediment trapping: obtained data



1. $TR = m/(Ft)$ - «trapping rate»' TR ($\text{gm}^{-2} \text{s}^{-1}$ or $\text{gm}^{-2}\text{d}^{-1}$)
2. Sedimentation rate for stream reach with bottom area $F1=F*n$:
 $SED = TR*F1$
3. Dimensionless relative settling coefficient : $C = TR/(vs)$

“Input” data: S – suspended sediment concentration, V – stream velocity, h – depth

Obtained data

25 field experiments in the different rivers



Protva river(

Koryak platue river

Oka river

Selenga river

Variation of trapping rate's for different rivers recorded by the sediment trap at different times (data bank fragment)

River	Distance from the bank, m	Depth h , m	Height above channel bed l , m	Relative depth $h = l/h$	Stream velocity v , m/s	Reinolds number $Re \cdot 10^{-3}$	SSC, g/m ³	Weight of material in each trap m , g	Average trapping rate C_{cp} , g·sm ⁻² ·s ⁻¹	Relative trapping rate $10^7 C_{cp} / (vs)$
Protva	1,5	0,8	0,55	0,69	0,1	79	7,1	3,1	0,02	2,82
								2,91		
								2,75		
	0,4	1	0,75	0,75	0,35	346	5,3	2,92	0,016	0,86
								1,32		
								2,99		
	1,5	0,5	0,25	0,50	0,15	74	4,4	2,61	0,022	3,33
								4,06		
								3,08		
	2,5	0,8	0,55	0,69	0,36	285	4,5	2,64	0,016	0,99
								1,66		
								2,61		
Koryak plateu rivers	1,5	0,35	0,10	0,29	0,15	52	1,44	0,005	0,00004	0,000048
	2,5	0,4	0,15	0,38	0,42	166	15	0,31	0,0017	0,0002
	2,5		0,06							
	2,0	0,3	0,05	0,17	0,49	145	4,3	0,17	0,004	0,0018
	2,5	0,4	0,15	0,38	0,45	178	8,0	0,11	0,0018	0,0005

Discussion: driving forces

gravitational vertical sediment fluxes

+

downstream movement of sediment particles

+

development of turbulent vortexes

Quantitative parameters

V

+

h_k/h

+

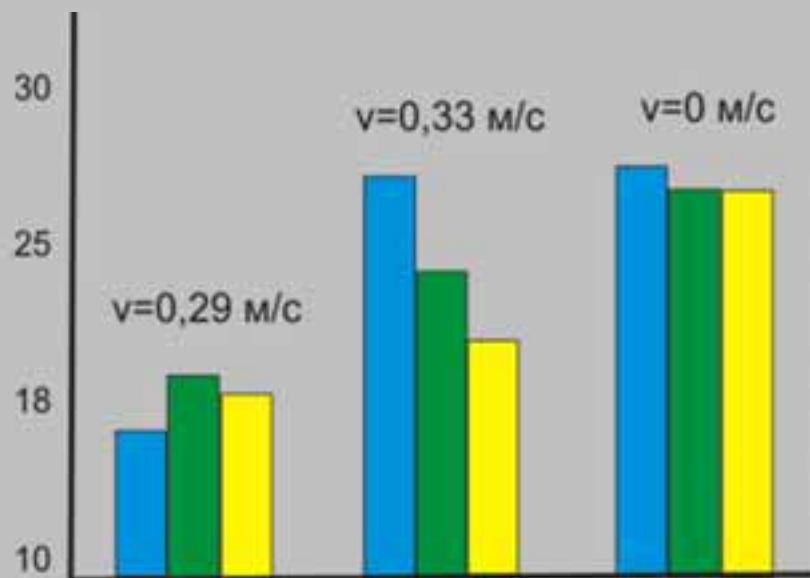
SSC

+

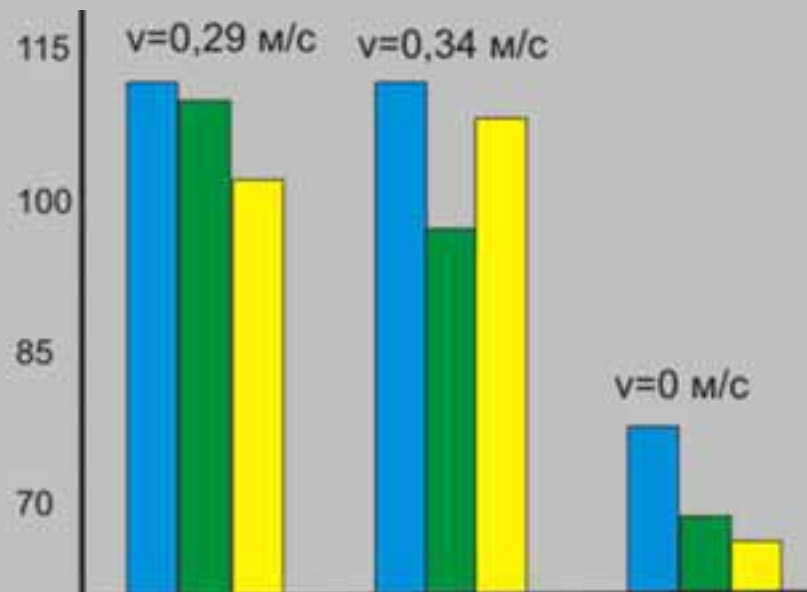
Re

Sedimentation at different flow velocities (river Protva)

TR, kg/m²sec



TR, kg/m²sec



- - high cylinder
- - middle cylinder
- - low cylinder

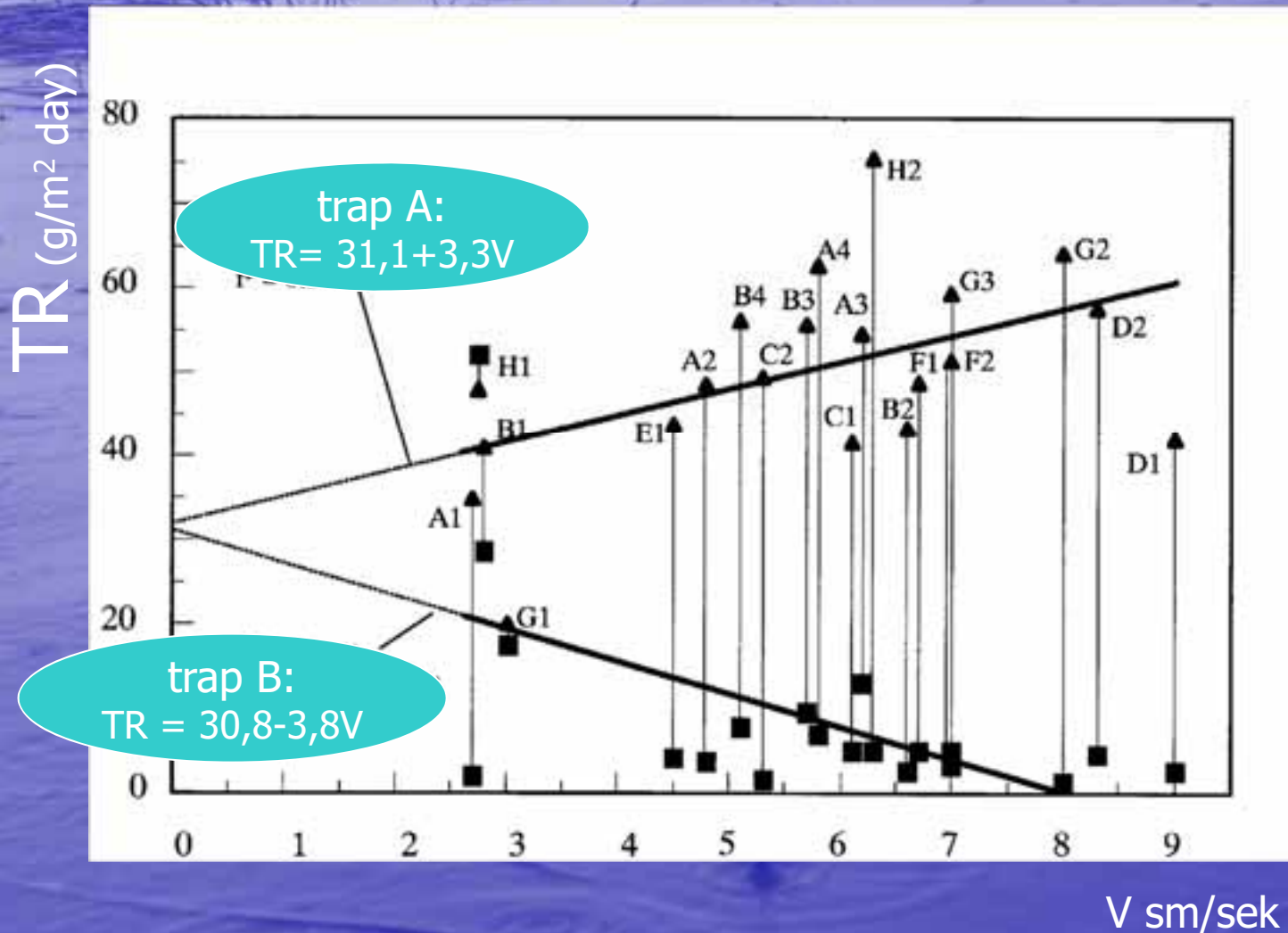


Low water

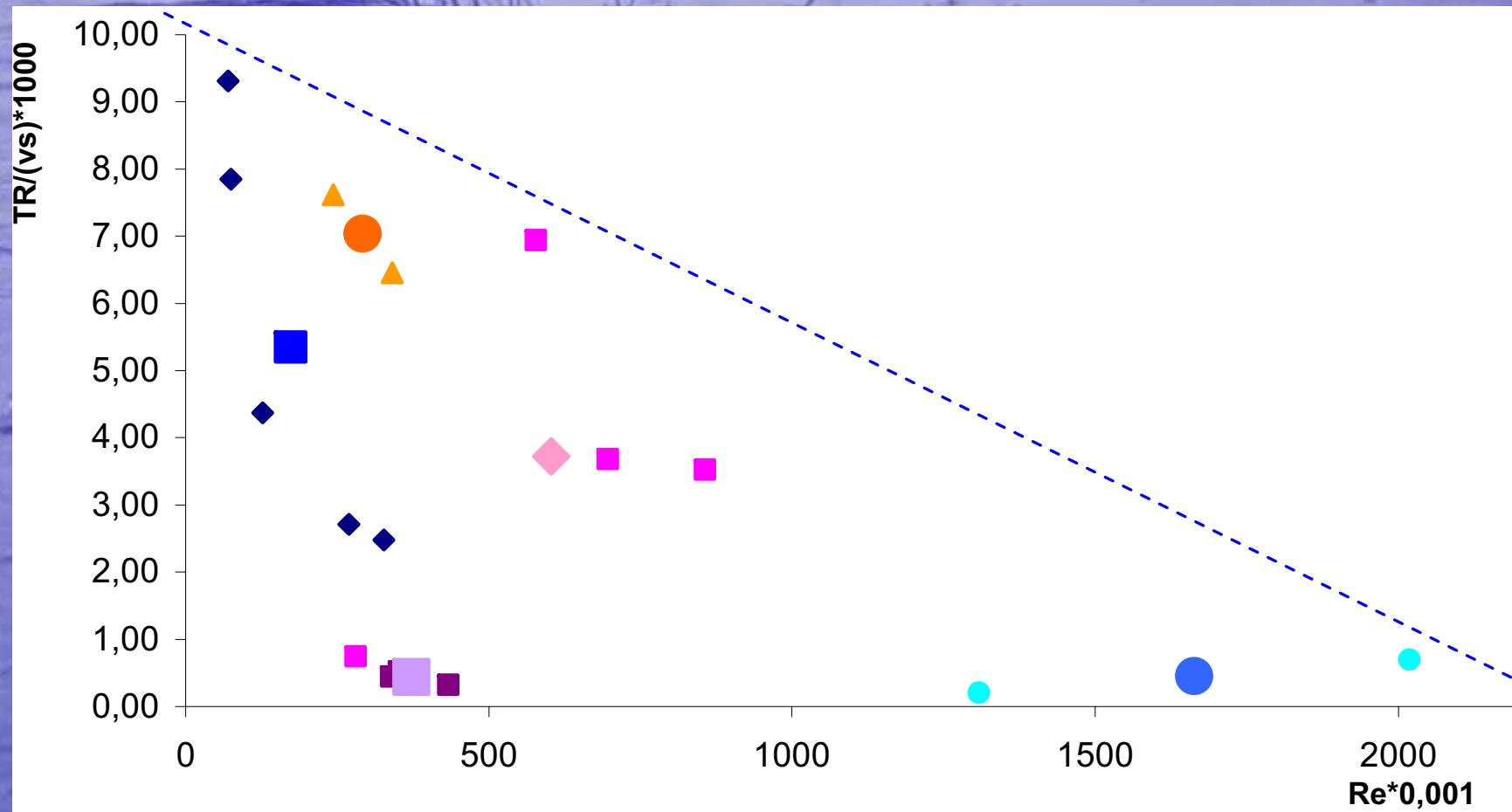
High water



Correlation between flow velocity v and «trapping rate» TR (by Kozerski (2006))



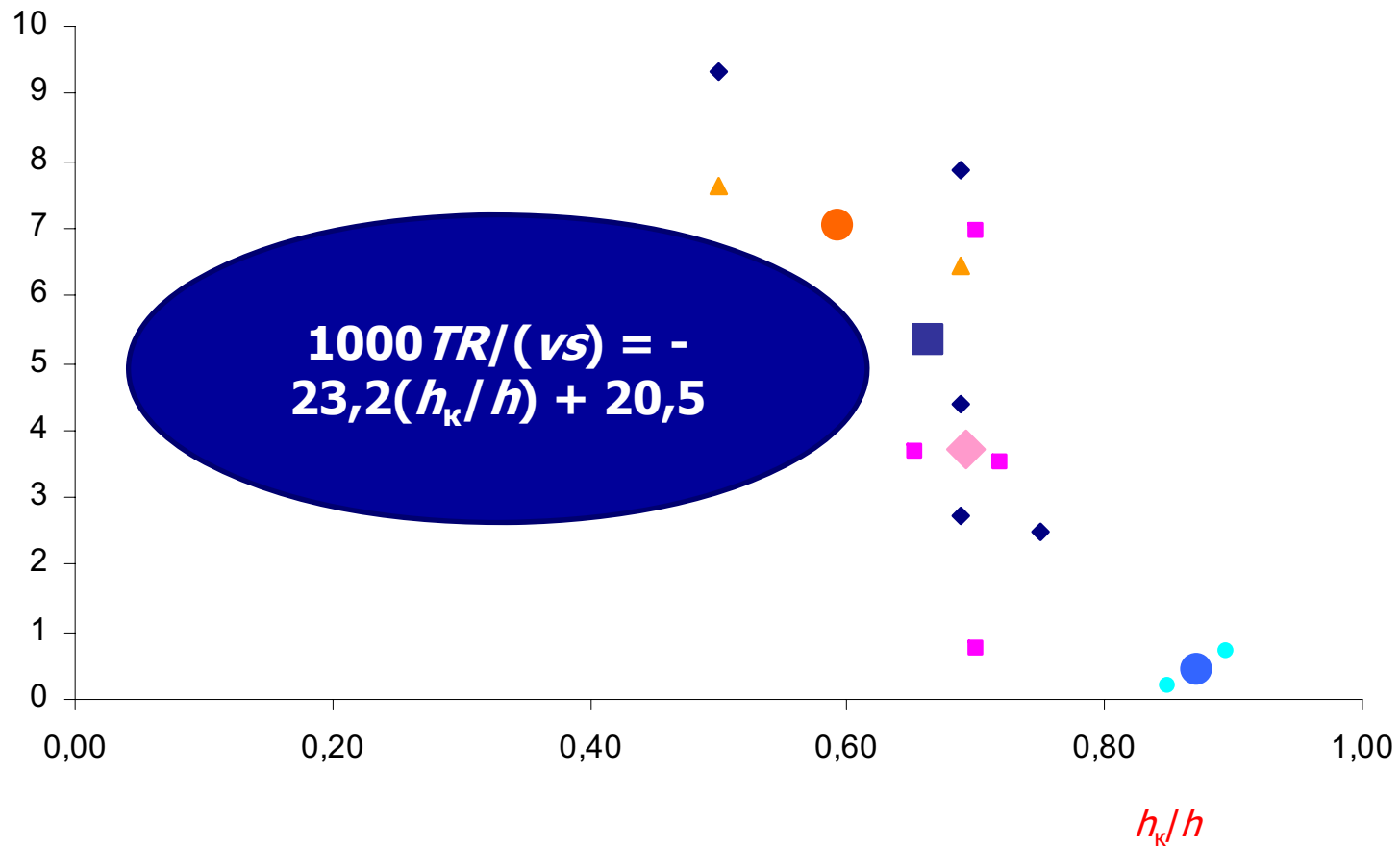
Relationship between Reynolds number and relative settling coefficient from trap installations on various rivers (marked by various symbols)



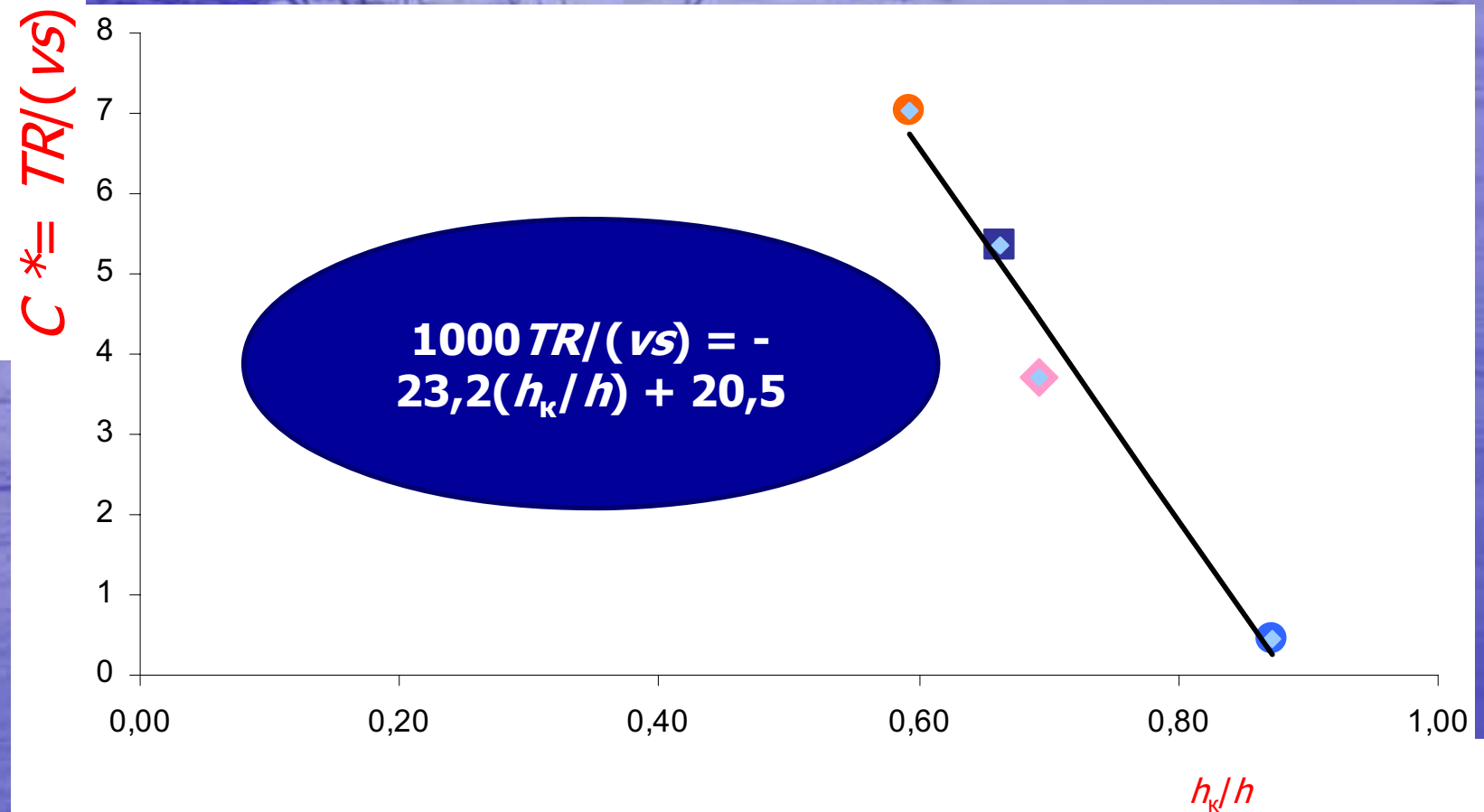
Dotted line – upper curve (corresponded to the highest sedimentation under $Re = \text{const}$)

Correlation between relative sediment trapping C^* and relative depth h_k/h (different size rivers)

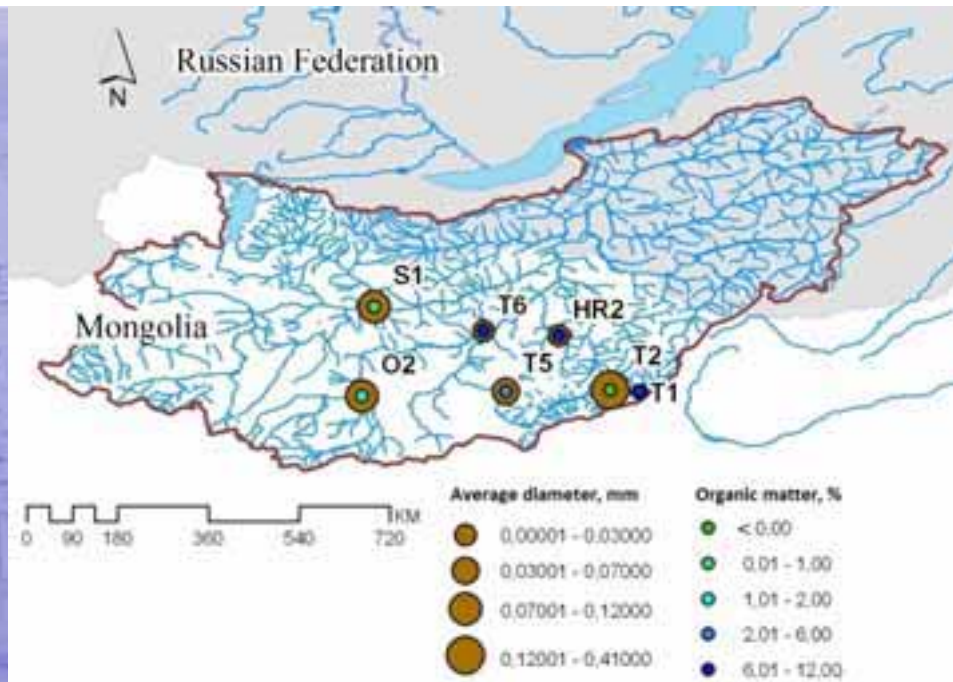
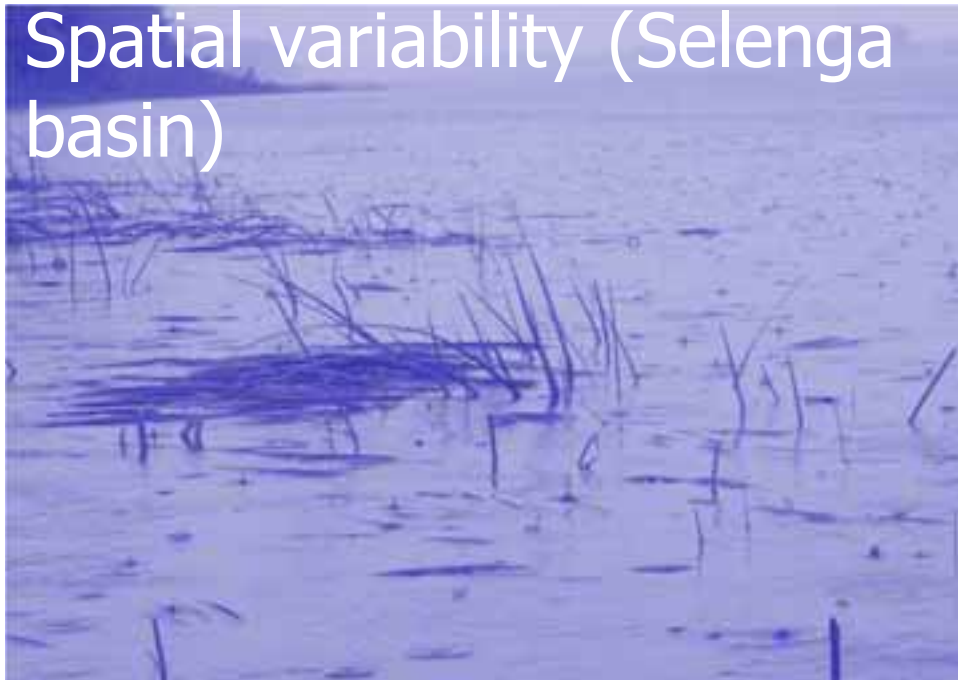
$C^* = TR/(vs)$



Correlation between relative sediment trapping C^* and relative depth h_k/h (different size rivers)



Spatial variability (Selenga basin)



Site	Description	SSC, g/m ³	Settling rate C, g/(m ² hour)	Suspended sediment load R, kg/s	Vertical sediment flux R _v , kg/s	R _v /R, %	Human impact
T1	Tuul river 20 km upstream from Ulaanbaatar	1.68	6.38	0.050	0.00003	0.06	low
T2	Tuul river 1 km upstream from Ulaanbaatar	8.36	87.5	0.22	0.00028	0.13	average
T5	Tuul river upstream from Zaamar	107	2048	3.56	0.027	0.75	average
T6	Tuul river downstream from Zaamar	289	1364	8.22	0.017	0.21	high
O2	Orkhon river upstream from the confluence with Tamir river	1699	2675	204	0.102	0.05	average
S1	Selenga river near Hutyk village	114	465	20.0	0.026	0.13	low

Comparison of sedimentation values obtained by trapping and direct measurements

Placer-mined rivers in the Kamchatka peninsula



Direct measuring thickness of fine sediments layer at stream bottom



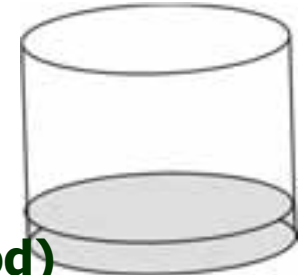
Obtained values of fine sediments layer thickness at the bottom

3-5 mm

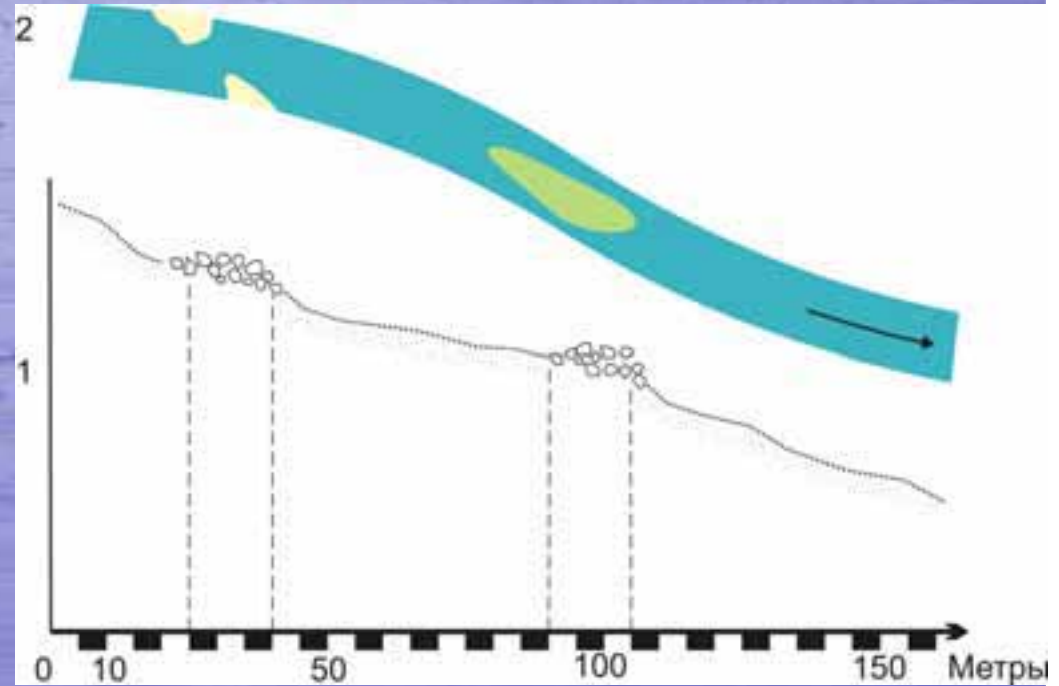
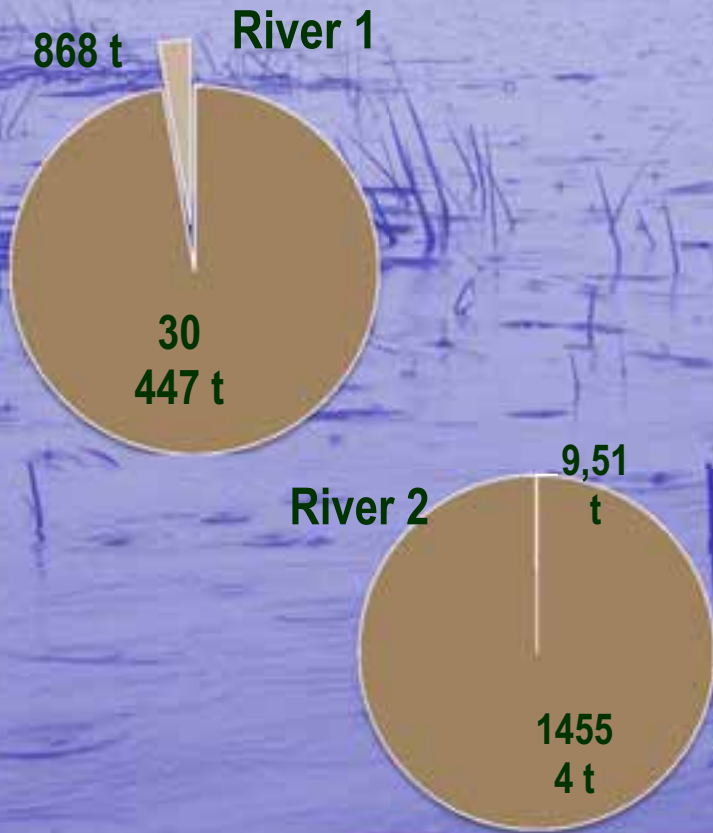
6 mm



Using point trap measurements TR (according to the average density of fine material, duration after high water period)



Contribution to sediment fluxes



River	Relative settling rates, kg/m ²		Settled particles layer, m		Volume of settled particles, m ³		MAss, t	
	Low water period duraton, days							
	30	75	30	75	30	75	30	75
1	1,98	4,94	0,0025	0,0062	434	1085	347	868
2	1,16	2,91	0,0015	0,0036	4,76	11,9	3,80	9,51

Conclusion

- Any attempt to provide a definitive assessment of the sedimentation processes in rivers causes faces a number of important constraints. Firstly it is instrumental error of trapping techniques which leads to the few apparent inconsistencies of the results obtained. Nevertheless absence of the analogue methods determines perspectives of the sediment trapping. It is regarded to be the only way in in-situ direct measurements of sedimentation.
- Future experiments will lead to the expanding of the data. Further multistatistical analyses will lead to describe exact laws and factors of sedimentation. Further investigations to introduce a method for the estimation of the average areal sedimentation flux for a section of a river based upon a series of point trap measurements should be done.

A photograph of a pond with lily pads and surrounding greenery. The pond is the central focus, filled with numerous lily pads. The water is calm, reflecting the sky and the surrounding trees. The background is a dense line of green trees under a bright sky. The foreground shows some tall grasses on the left and right sides of the pond.

**Thank you for your
attention!**

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